The hydrostatic approximation for stratified fluids

Vincent Duchêne

CNRS & IRMAR, Univ. Rennes 1

Programme CEA-SMAI/GAMNI Institut Henri Poincaré, Paris, janvier 2022

Joint work with Roberta Bianchini (CNR, Rome)

Overview



2 The hydrostatic problem

3 Eddy diffusivity

- Isopycnal coordinates
- 5 The hydrostatic limit



Motivated by stratified flows, we study the inhomogeneous, incompressible Euler equations with gravity force:

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \rho (\partial_t w + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_z w) + \partial_z P + g \rho = 0, \\ \nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_z w = 0, \\ \text{boundary conditions.} \end{cases}$$
(E)

The velocity field is denoted $(\mathbf{u}, w) : \Omega \to \mathbb{R}^d \times \mathbb{R}$, the density $\rho : \Omega \to \mathbb{R}^+_*$. The pressure $P : \Omega \to \mathbb{R}$ satisfies the elliptic equations

 $-\nabla_{\mathbf{x}} \cdot \left(\frac{1}{\rho} \nabla_{\mathbf{x}} P\right) - \partial_{z} \left(\frac{1}{\rho} \partial_{z} P\right) = \nabla_{\mathbf{x}} \cdot \left((\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_{z} \mathbf{u} \right) + \partial_{z} \left(g \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_{z} w \right).$

After rescaling $\mathbf{x} \leftarrow L_{\mathbf{x}}\mathbf{x}$, $z \leftarrow L_{z}z$, (\cdots) , the system becomes ...



Motivated by stratified flows, we study the inhomogeneous, incompressible Euler equations with gravity force:

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \rho (\partial_t w + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_z w) + \partial_z P + g \rho = 0, \\ \nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_z w = 0, \\ \text{boundary conditions.} \end{cases}$$
(E)

The velocity field is denoted $(\mathbf{u}, w) : \Omega \to \mathbb{R}^d \times \mathbb{R}$, the density $\rho : \Omega \to \mathbb{R}^+_*$. The pressure $P : \Omega \to \mathbb{R}$ satisfies the elliptic equations

 $-\nabla_{\mathbf{x}} \cdot \left(\frac{1}{\rho} \nabla_{\mathbf{x}} P\right) - \partial_{z} \left(\frac{1}{\rho} \partial_{z} P\right) = \nabla_{\mathbf{x}} \cdot \left((\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_{z} \mathbf{u} \right) + \partial_{z} \left(g \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_{z} w \right).$

After rescaling $\mathbf{x} \leftarrow L_{\mathbf{x}}\mathbf{x}$, $z \leftarrow L_{z}z$, (\cdots) , the system becomes ...



The initial set of equations

Motivated by stratified flows, we study the inhomogeneous, incompressible Euler equations with gravity force:

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \frac{L_z^2}{L_x^2} \rho (\partial_t w + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_z w) + \partial_z P + \rho = 0, \\ \nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_z w = 0, \\ \text{boundary conditions.} \end{cases}$$
(E)

The velocity field is denoted $(\mathbf{u}, w) : \Omega \to \mathbb{R}^d \times \mathbb{R}$, the density $\rho : \Omega \to \mathbb{R}^+_*$. The pressure $P : \Omega \to \mathbb{R}$ satisfies the elliptic equations

$$-\nabla_{\mathbf{x}} \cdot \left(\frac{1}{\rho} \nabla_{\mathbf{x}} P\right) - \frac{L_{\mathbf{x}}^2}{L_{\mathbf{z}}^2} \partial_{\mathbf{z}} \left(\frac{1}{\rho} \partial_{\mathbf{z}} P\right) = \nabla_{\mathbf{x}} \cdot \left(\left(\mathbf{u} \cdot \nabla_{\mathbf{x}}\right) \mathbf{u} + w \partial_{\mathbf{z}} \mathbf{u} \right) + \partial_{\mathbf{z}} \left(\frac{L_{\mathbf{x}}^2}{L_{\mathbf{z}}^2} \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_{\mathbf{z}} w \right).$$

After rescaling $\mathbf{x} \leftarrow L_{\mathbf{x}}\mathbf{x}$, $z \leftarrow L_{z}z$, (\cdots) , the system becomes ...



The hydrostatic approximation

Formally, the Euler equations

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \frac{L_z^2}{L_x^2} \rho (\partial_t w + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_z w) + \partial_z P + \rho = 0, \\ \nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_z w = 0, \end{cases}$$
(E) boundary conditions.

in the <u>hydrostatic limit</u> $\mu \stackrel{\text{def}}{=} \frac{L_z^2}{L_x^2} \to 0$, read $\begin{cases}
\partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\
\rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\
\partial_z P = -\rho, \\
\partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\
\text{boundary conditions.}
\end{cases}$ $\implies P = P|_{z=z_{\text{surf}}} + \int_z^{z_{\text{surf}}} \rho \, \mathrm{d}z', \quad w = w|_{z=z_{\text{bot}}} - \int_{z_{\text{bot}}}^z \nabla_{\mathbf{x}} \cdot \mathbf{u} \, \mathrm{d}z'.$ (H)

Context The hydrostatic problem Eddy diffusivity Isopycnal coordinates The hydrostatic I 00 00 00 00 00 00 00 (partial and biased) state of the art on the hydrostatic equations hydrostatic equations 00 00 00

Homogeneous case: $\rho \equiv 1$.

- Spectral stability of the linearized system about shear flows, under the Rayleigh criterion $\underline{\mathbf{u}}''(z) \neq 0$. [Rayleigh (1880)][Arnold][Drazin&Reid]
- Ill-posedness of the linearized system about certain shear flows. [Renardy '09]
- Well-posedness of the (nonlinear) system in Sobolev spaces. [Masmoudi&Wong '12] (after [Grenier '99], [Brenier '03])

Stably stratified case: $\partial_z \rho < 0$.

• Spectral stability of the linearized system about shear flows, under the Miles and Howard criterion $|\underline{\mathbf{u}}'(z)|^2 \leq \frac{1}{4} \frac{-\rho'(z)}{\rho(z)}$. [Miles '61][Howard '61]

•

Context OOO
The hydrostatic problem OO
Eddy diffusivity OO
Eddy diffu

Let us linearize the system

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \\ \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (periodic).} \end{cases}$$

about shear solutions $(\rho, \mathbf{u})(t, \mathbf{x}, z) = (\underline{\rho}(z), \underline{\mathbf{u}}(z))$. We get

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho + \underline{\rho}'(z) w = 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + \underline{\mathbf{u}}'(z) w + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \qquad (\Longrightarrow P = \int_z^0 \rho \ dz' \stackrel{\text{def}}{=} \mathsf{L}\rho) \\ \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \qquad (\Longrightarrow w = -\int_{-1}^z \nabla_{\mathbf{x}} \cdot \mathbf{u} \ dz') \\ \text{boundary conditions (periodic).} \end{cases}$$
(L)

(H)

The hydrostatic problem • The linearized system about shear flows

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho + \underline{\rho}'(z) w = 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + \underline{\mathbf{u}}'(z) w + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \qquad (\Longrightarrow P = \int_z^0 \rho \, \mathrm{d}z' \stackrel{\mathrm{def}}{=} \mathsf{L}\rho) \\ \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \qquad (\Longrightarrow w = -\int_{-1}^z \nabla_{\mathbf{x}} \cdot \mathbf{u} \, \mathrm{d}z') \\ \text{boundary conditions (periodic).} \end{cases}$$
(L

The linearized system enjoys a partial symmetric structure

$$\begin{aligned} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} &= 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho &= 0, \end{aligned} \tag{L}$$
 boundary conditions (periodic).

$$\frac{\mathrm{d}}{\mathrm{d}t} \left(\left(\rho , \frac{-1}{\underline{\rho}'(z)} \rho \right)_{L^2_{\mathbf{x},z}} + \left(\mathbf{u} , \underline{\rho}(z) \mathbf{u} \right)_{L^2_{\mathbf{x},z}} \right) = \left(\mathbf{u} , \underline{\rho}(z) \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} \right)_{L^2_{\mathbf{x},z}}.$$

 \rightsquigarrow no (obvious) control of the energy.

Context 000	The hydrostatic problem ⊙●	Eddy diffusivity 00	lsopycnal coordinates	The hydrostatic limit
		Conclusion		

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} = 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho = 0, \\ \text{boundary conditions (periodic).} \end{cases}$$
(L)

 $\frac{\mathrm{d}}{\mathrm{d}t} \left(\left(\rho , \frac{-1}{\underline{\rho}'(z)} \rho \right)_{L^{2}_{\mathbf{x},z}} + \left(\mathbf{u} , \underline{\rho}(z) \mathbf{u} \right)_{L^{2}_{\mathbf{x},z}} \right) = \left(\mathbf{u} , \underline{\rho}(z) \underline{\mathbf{u}}'(z) \mathsf{L}^{*} \nabla_{\mathbf{x}} \cdot \mathbf{u} \right)_{L^{2}_{\mathbf{x},z}}.$

No (obvious) control of the energy in the presence of shear velocity, $\underline{\mathbf{u}}'(z) \neq 0$, even under the Miles&Howard criterion. The stable stratification, $\rho'(z) < 0$, helps (a bit).

Possible ways around:

- Work in the analytic framework. [Kukavica, Temam, Vicol & Ziane '11]
- Add (horizontal) viscosity. [Cao, Li & Titi '16]
- Do both. [Paicu, Zhang& Zhang '20]

ext The hydrostatic problem

Eddy diffusivity

Isopycnal coordinates

The hydrostatic limit

Eddy viscosity and diffusivity

The system studied in [Cao, Li & Titi '16] is

$$\begin{aligned} \partial_t \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} \cdot \rho + w \partial_z \rho &= \kappa \Delta_{\mathbf{x}} \rho, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P &= \nu \Delta_{\mathbf{x}} \mathbf{u}, \\ \partial_z P &= -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (periodic).} \end{aligned}$$
(Figure 1.1)

The <u>horizontal</u> viscosity ($\nu > 0$) and diffusivity ($\kappa > 0$) approximate effective isopycnal viscosity and diffusivity:

[Gent & McWilliams '90] propose

with

$$\begin{cases} \partial_t \rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}} \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t \mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + (w + w_{\star}) \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions.} \end{cases}$$
(H)

$$\mathbf{u}_{\star} = \kappa \partial_{\mathbf{z}} \left(\frac{\nabla_{\mathbf{x}} \rho}{\partial_{\mathbf{z}} \rho} \right) \ , \quad \mathbf{w}_{\star} = -\kappa \nabla_{\mathbf{x}} \cdot \left(\frac{\nabla_{\mathbf{x}} \rho}{\partial_{\mathbf{z}} \rho} \right) \ .$$

Context	The hydrostatic problem	Eddy diffusivity	Isopycnal co
000	00	0.	00

The hydrostatic limit

Main result

$$\begin{cases} \partial_t \rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}} \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t \mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + (w + w_{\star}) \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (free surface).} \end{cases}$$
(H)

[VD & R. Bianchini]

For sufficiently regular data satisfying the (strict) stable stratification assumption

 $-\partial_z \rho|_{t=0} \geq \alpha > 0$

and for any $\kappa > 0$, there exists a unique (classical) solution of (H) on the time interval [0, T] with

$$\mathcal{T}^{-1} = \mathcal{C} \left(1 + \kappa^{-1} \left(\left| \underline{\mathbf{u}}' \right|_{L^2_{\tau}}^2 + M_0^2 \right) \right),$$

where M_0 is the size of the initial deviation from the shear flow equilibrium $(\rho(z), \underline{\mathbf{u}}(z))$, and C depends only on M_0 , α and the size of $(\rho(z), \underline{\mathbf{u}}(z))$.

ntext	The	hydrostatic	problem
0	00		

Eddy diffusivity O• sopycnal coordinates

The hydrostatic limit

[VD & R. Bianchini]

For sufficiently regular data satisfying the (strict) stable stratification assumption

Main result

 $-\partial_z \rho|_{t=0} \geq \alpha > 0$

and for any $\kappa > 0$, there exists a unique (classical) solution of (H) on the time interval [0, T] with

$$T^{-1} = C \left(1 + \kappa^{-1} \left(\left| \underline{\mathbf{u}}' \right|_{L^{2}_{z}}^{2} + M_{0}^{2} \right) \right),$$

where M_0 is the size of the initial deviation from the shear flow equilibrium $(\rho(z), \underline{\mathbf{u}}(z))$, and C depends only on M_0 , α and the size of $(\rho(z), \underline{\mathbf{u}}(z))$.

Key points.

- We do not use viscosity, nor analytic data;
- The destabilizing role of the shear velocity is apparent;
- Stable stratification is used in a crucial way.



We define the variable $h(t, \mathbf{x}, \varrho) > 0$ from the density $\rho(t, \mathbf{x}, z)$ through

$$h \stackrel{\text{def}}{=} -\partial_{\varrho}\zeta, \qquad \rho(t, \mathbf{x}, \zeta(t, \mathbf{x}, \varrho)) = \varrho, \quad \zeta(t, \mathbf{x}, \rho(t, \mathbf{x}, z)) = z.$$



The hydrostatic problemEddy diffusivit0000

Isopycnal coordinates

The hydrostatic limit

Isopycnal coordinates

We define the variable $h(t, \mathbf{x}, \varrho) > 0$ from the density $\rho(t, \mathbf{x}, z)$ through

$$h \stackrel{\mathrm{def}}{=} -\partial_{\varrho}\zeta, \qquad
ho(t, \mathbf{x}, \zeta(t, \mathbf{x}, \varrho)) = \varrho, \quad \zeta(t, \mathbf{x}, \rho(t, \mathbf{x}, z)) = z.$$

The system

$$\begin{cases} \partial_t \rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}} \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t \mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + (w + w_{\star}) \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (free surface).} \end{cases}$$
(H)

reads in isopycnal coordinates

$$\begin{cases} \partial_t h + \nabla_{\mathbf{x}} \cdot (h\mathbf{u}) = \kappa \Delta_{\mathbf{x}} h, \\ \varrho \Big(\partial_t \mathbf{u} + \big((\mathbf{u} + \kappa \frac{-\nabla_{\mathbf{x}} h}{h}) \cdot \nabla_{\mathbf{x}} \big) \mathbf{u} \Big) + \nabla_{\mathbf{x}} \psi = 0, \end{cases}$$
(H)

where

$$\psi(t,\mathbf{x},\varrho) \stackrel{\text{def}}{=} \rho_0 \int_{\rho_0}^{\rho_1} h(t,\mathbf{x},\varrho') \, \mathrm{d}\varrho' + \int_{\rho_0}^{\varrho} \int_{\varrho'}^{\rho_1} h(t,\mathbf{x},\varrho'') \, \mathrm{d}\varrho'' \, \mathrm{d}\varrho',$$

Context	The hydrostatic problem	Eddy diffusivity	Isopycnal coordinates	The hydrostatic limit
000	00	00	0.	000
	TTL	• •	1	

The system in isopycnal coordinates

$$\begin{cases} \partial_t h + \nabla_{\mathbf{x}} \cdot (h\mathbf{u}) = \kappa \Delta_{\mathbf{x}} h, \\ \varrho \Big(\partial_t \mathbf{u} + \Big((\mathbf{u} + \kappa \frac{-\nabla_{\mathbf{x}} h}{h} \Big) \cdot \nabla_{\mathbf{x}} \Big) \mathbf{u} \Big) + \nabla_{\mathbf{x}} \psi = 0. \end{cases}$$
(H)

Key points:

- The advection in the variable *z* has disappeared (semi-Lagrangian coordinates).
- The domain is flattened.
- The system is easily discretized (multilayer framework).
- The eddy viscosity of [Gent & McWilliams '90] is nice and simple.
- The system enjoys a partial symmetric structure analogous as the one exhibited in the Eulerian coordinates, but the symmetry defect involves a derivative on *h* (not **u**).

 \rightsquigarrow The proof of the results is based on the energy method on (H), using product, commutator, composition estimates in anisotropic Sobolev spaces.



The standard procedure to prove convergence (with a convergence rate, and loss of derivatives) of regular solutions of the Euler system towards corresponding solutions of the hydrostatic system as $\mu \stackrel{\text{def}}{=} \frac{L_z^2}{L_x^2} \rightarrow 0$, rely on with

- the existence and control of solutions to the limit hydrostatic system;
- a consistency result on the latter solutions;
- the (small-time) existence of solutions to the Euler equations,
- uniform stability estimates on the linearized Euler equations.

Remark: in [Desjardins, Lannes & Saut '20], the authors show the well-posedness on small (that is non-uniform) time of the Euler system (without diffusivity, and rigid-lid assumption).

Eddy diffusivity Isopycnal coordinates 00 00 00

The hydrostatic limit

(E)

The non-hydrostatic problem

In isopycnal coordinates, the Euler equations read

$$\begin{cases} \partial_t h + \nabla_{\mathbf{x}} \cdot (h\mathbf{u}) = \kappa \Delta_{\mathbf{x}} h, \\ \varrho \Big(\partial_t \mathbf{u} + \big(\big(\mathbf{u} - \kappa \frac{\nabla_{\mathbf{x}} h}{h} \big) \cdot \nabla_{\mathbf{x}} \big) \mathbf{u} \Big) + \nabla_{\mathbf{x}} P + \frac{\nabla_{\mathbf{x}} \zeta}{h} \partial_{\varrho} P = 0, \\ \mu \varrho \Big(\partial_t w + \big(\mathbf{u} - \kappa \frac{\nabla_{\mathbf{x}} h}{h} \big) \cdot \nabla_{\mathbf{x}} w \Big) - \frac{\partial_{\varrho} P}{h} + \varrho = 0, \\ -h \nabla_{\mathbf{x}} \cdot \mathbf{u} - \nabla_{\mathbf{x}} \zeta \cdot \partial_{\varrho} \mathbf{u} + \partial_{\varrho} w = 0, \quad \text{(incomp.)} \\ P \Big|_{\varrho = \rho_0} = 0, \quad w \Big|_{\varrho = \rho_1} = 0, \quad \text{(bound. cond.)} \end{cases}$$

where $\zeta(\cdot, \varrho) \stackrel{\text{def}}{=} \int_{\varrho}^{\rho_1} h(\cdot, \varrho') \, \mathrm{d}\varrho'$.

The pressure reconstruction system becomes

$$\begin{pmatrix} \frac{1}{\mu} \begin{pmatrix} \sqrt{\mu} \nabla_{\mathbf{x}} \\ \partial_{\varrho} \end{pmatrix} \cdot \left(\begin{pmatrix} \frac{h}{\varrho} \operatorname{Id} & \frac{\sqrt{\mu} \nabla_{\mathbf{x}} \zeta}{\varrho} \\ \frac{\sqrt{\mu} \nabla_{\mathbf{x}}^{\top} \zeta}{\varrho} & \frac{1+\mu |\nabla_{\mathbf{x}} \zeta|^2}{\varrho h} \end{pmatrix} \begin{pmatrix} \sqrt{\mu} \nabla_{\mathbf{x}} \\ \partial_{\varrho} \end{pmatrix} P \right) = \operatorname{RHS}$$

$$\left. \begin{pmatrix} P |_{\varrho=\rho_0} = 0, & (\partial_{\varrho} P) |_{\varrho=\rho_1} = \rho_1 h |_{\varrho=\rho_1}. \end{cases}$$

where RHS is long, ugly, but nice.

The hydrostatic problem

Eddy diffusivity

Isopycnal coordinates

The hydrostatic limit

Symmetric structure

Decomposing $P = P_{hydro} + \tilde{P}$, (with $P_{hydro}(\cdot, \varrho) := \int_{\rho_0}^{\varrho} \varrho' h(\cdot, \varrho') d\varrho'$) we have $\begin{cases} \partial_t h + \nabla_{\mathbf{x}} \cdot (h\mathbf{u}) = \kappa \Delta_{\mathbf{x}} h, \\ \varrho \Big(\partial_t \mathbf{u} + \big((\mathbf{u} - \kappa \frac{\nabla_{\mathbf{x}} h}{h} \big) \cdot \nabla_{\mathbf{x}} \big) \mathbf{u} \Big) + \nabla_{\mathbf{x}} \psi_{hyd.} + \nabla_{\mathbf{x}} \tilde{P} + \frac{\nabla_{\mathbf{x}} \zeta}{h} \partial_{\varrho} \tilde{P} = 0, \\ \mu \varrho \Big(\partial_t w + \big(\mathbf{u} - \kappa \frac{\nabla_{\mathbf{x}} h}{h} \big) \cdot \nabla_{\mathbf{x}} w \Big) - \frac{\partial_{\varrho} \tilde{P}}{h} = 0, \quad (E) \\ -h \nabla_{\mathbf{x}} \cdot \mathbf{u} - \nabla_{\mathbf{x}} \zeta \cdot \partial_{\varrho} \mathbf{u} + \partial_{\varrho} w = 0, \quad (incomp.) \\ \tilde{P} \Big|_{\varrho = \rho_0} = 0, \quad w \Big|_{\varrho = \rho_1} = 0, \quad (bound. \text{ cond.}) \end{cases}$

where $(\sqrt{\mu}\nabla_{\mathbf{x}}\tilde{P},\partial_{\varrho}\tilde{P})$ is of size $\mathcal{O}(1)$ without loss of derivatives, $\mathcal{O}(\mu)$ with loss.

We observe a symmetric structure superposed to the one of the hydrostatic system, which provides the desired estimates.

[VD & R. Bianchini]

For any sufficiently regular solution satisfying the stable stratification assumption $h \ge \alpha > 0$, and for any $\kappa > 0$, we have existence and uniqueness of solutions to the Euler equation and strong convergence towards the hydrostatic solution as

 $\mu \rightarrow 0.$

Thoughts to go

- The well-posedness of the non-homogeneous hydrostatic equations, without diffusivity or viscosity, is an open problem.
- Stable stratification helps, but seemingly (?) not enough.
- Isopycnal coordinates are quite interesting for numerical and theoretical analyses.

Thoughts to go

- The well-posedness of the non-homogeneous hydrostatic equations, without diffusivity or viscosity, is an open problem.
- Stable stratification helps, but seemingly (?) not enough.
- Isopycnal coordinates are quite interesting for numerical and theoretical analyses.

Thank you for your attention