# Earthquakes and water

(and why the LUSI eruption was not triggered by an Earthquake)



Michael Manga, University of California, Berkeley

Interrupt and ask questions at any time

# Why?

"It's one of those curiosities of nature that has preoccupied people for years" Stuart Rojstaczer, *Nature*, 2003

Hope to obtain new insight into interactions between stress and subsurface fluid flow, transport and the evolution of hydrogeologic properties

Long history of study (e.g., Pliny, 77 AD) but still not understood

#### Formerly dry stream that starts to flow after the earthquake (video by Mike Hendry of Hendry winery)



San Francisco Chronicle, Sunday Sept 8, 2014 cover story: "Surprise bonanza since Napa quake: dry creeks now flowing"

# Earthquakes and water

(and why the LUSI eruption was not triggered by an Earthquake)

Examples of hydrological responses to earthquakes

- Liquefaction
- Mud volcanoes (and magmatic ones)
- Stream flow
- Wells (level, temperature, geochemistry)
- Geysers
- Earthquakes
- Precursors

# Why does streamflow increase after (distant) earthquakes?



Observed and documented for thousands of years

Distant: more than 1 fault length away from ruptured fault

### Origin of excess flow $= -AK\partial h/\partial x$ Discharge Head Hydraulic gradient conductivity

 Hydraulic conductivity (permeability) increased
 Head gradient increased (water released from storage)

# Five hypotheses

1) Static compression (e.g., Muir-Wood and King, 1993)

2) Breaching barriers or seals (e.g., Sibson 1994; Brodsky et al. 2003)

3) Consolidation and liquefaction

(e.g., Manga, 2001; Montgomery et al., 2003)

4) Shaken from the unsaturated zone (Mohr, Manga et al., 2015)

#### 5) Permeability enhancement

(e.g., Rojstaczer et al., 1995; Tokunaga, 1999; Sato et al., 2000)

Static stress: permanent changes in stress caused by displacement along the fault

Dynamic stress: temporary stresses created by passage of seismic waves

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#### Why consolidation or liquefaction?

Similar magnitude-distance relationships



Where is the excess water coming from?

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# New streams and springs after the 2014 M6 South Napa earthquake



Key: streams were dry before earthquake No significant rain for ~ 2 months after earthquake

#### Geologic setting

The Napa-Sonoma Valleys are structural troughs between the Coast Ranges north of San Francisco Bay

Valleys are filled with Quaternary alluvial deposits that overlie Plio-Pleistocene sedimentary rocks, which in turn overlie Miocene to Pliocene Sonoma volcanics of uncertain thickness.

Mountains consist mostly of the Sonoma volcanics, with soil cover rarely exceeding a meter in thickness.



#### 0. Where is water coming from?



Water emerges at base of mountains, not in the alluvial deposits



Flow ALWAYS increased, even for small or negative volume strain (implication, if correct: given a few mm of new flow and strains, water must be expelled from >10 km of crust!)

#### 2. New (hydrothermal) water?

Concentrations of B and Li in the new streams are much more similar to the regional groundwater than to hydrothermal fluids

#### 3. Liquefaction?

Only a few instances of ground failure or liquefaction (Brocher et al., 2015) and excess water is from the mountains

#### 4. Unsaturated zone?

Sampled waters for O and H isotopes, and find no signature of evaporation in new flows

details and data in Wang and Manga, Nature Communications (2015)

#### 2. New (hydrothermal) water?

Sample	Date	Li (ppb)	B (ppb)
Oakville 1	10/1/14	27.1	431
	10/11/14	27.8	447
Oakville 2	10/1/14	32.6	587
	10/11/14	38.1	467
Oakville 3	10/1/14	9.7	537
	10/11/14	27.5	574
Felder	10/1/14	27.7	92
	10/11/14	28.4	96
Апоуо Seco	10/1/14	8.3	43
-	10/11/14	9.1	48
Carriger	10/1/14	3.6	18
	10/11/14	3.7	19
Nathanson	10/1/14	1.8	18
	10/11/14	1.8	18
Spencer Spring	10/11/14	79.4	739
*Hydrothermal fluid		1633	10723
*Groundwater		15.5	211

\* Forrest et al., Applied Geochemistry (2013) Measurements by Ben Thurhoffer and Jim Bishop

Concentrations of B and Li in the new streams are much more similar to the regional groundwater than to hydrothermal fluids (except the warm spring)



Only a few instances of ground failure or liquefaction (Brocher et al., 2015) excess water from the mountains

#### 4. Unsaturated zone?



#### 4. Unsaturated zone?

Small seasonal changes in perennial streams; no strong evidence for an evaporation signal





Solve groundwater flow equations analytically

Fit model to data (obtain hydraulic diffusivity and total excess discharge)

Vertical permeability enhancement model of Wang et al., *Geology* (2004)







Total excess discharge  $\sim 10^6 \text{ m}^3$ 

Drainage of water would make crust more compressible and lower seismic velocity of shallow crust



Taira et al., *GRL* (2015) document a velocity decrease Not correlated with static strain

Drainage of water would make crust more compressible and lower seismic velocity of shallow crust



Taira et al., *GRL* (2015) document a velocity decrease There IS a correlation with dynamic strain

Drainage of water would make crust more compressible and lower seismic velocity of shallow crust



Taira et al., GRL (2015) document a velocity decrease Recovery time scales ~ time with excess discharge

## Conclusions

1) Static compression (e.g., Muir-Wood and King, 1993)

2) Breaching barriers or seals (e.g., Sibson 1994; Brodsky et al. 2003)

3) Consolidation and liquefaction (e.g., Manga, 2001; Montgomery et al., 2003)

4) Shaken from the unsaturated zone (Mohr et al., 2015)

5) Permeability enhancement by dynamic strain (e.g., Rojstaczer et al., 1995; Tokunaga, 1999; Sato et al., 2000)



# Why Chile? Strong climatic gradients

- High seismicity
- Steep topography gradients
- Diverse geology
- Long time series of

hydrometeorological data

Details in Mohr et al., EPSL (2017)

#### Data available:

- 716 catchments monitored (1940 -)
- 802 rainfall gauges (1940 -)
- 75 Air temperature stations (1960 -)

#### Earthquakes covered:

- M8.1 Coquimbo 1943
- M8.2 Antofagasta 1950
- M9.5 Valdivia 1960
- M8.0 Valparaiso 1985
- M8.0 Antofagasta 1995
- M8.8 Maule 2010
- M8.2 Iquique 2014



#### Streamflow responses to M8.8 Maule earthquake 2010





- Volumetric strain uncorrelated with sign of responses
- Flow decreased for positive volumetric strains
- Potential effects of hydropower excluded



https://file.ejatlas.org/docs/ralco\_chile2.jpg, 12/07/2015



2500 100 Change in streamflow (%) 50 2000 0 1500 -50 5×10° 1.5×105 1000 500 0 .5×10 5×10 1×10 5×10 2×10 0

Static strain

M8.8 Maule EQ

- Volumetric strain uncorrelated with sign of responses
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https://file.ejatlas.org/docs/ralco\_chile2.jpg, 12/07/2015

#### 3+5. Liquefaction from dynamic strain? Vadose zone?



- Seismic energy sufficient to trigger streamflow responses
- Undrained consolidation possible, but no field evidence
- Excess water also in mountains

Papadopoulos and Lefkopoulos (1993)

#### 5. Enhanced permeability?





But the model may not explain all the observations...




		0	5	10	15
G_prc95 G_prc01 G_prc01	is	•			
G_kurtos G_prc33	IS				
G_std G_mean		•			
G_prc05 G_max		•			
Z_mean G_skewr	ess	•			
C Z_media G_prc99	n	•			
G_media Z_prc95	n	•			
Lange C_range		•			
Z_std					
Z_prcss Z_max G_prc66			-		
Z_prc33 Z_prc66 Z_prc99					
Z_skewn Z_prc05	ess		•		
ice_glaci	ers	•			
unconso metamo	idated phic	• • • • • •			
volcanic catchme	nt area	•			
unconso closest_f	idated_volcanic ault idated_sedimentary		•		
Sedimen	nsity ary		•		
rainfall_a rainfall_f	nnual eb		•		
wetlands					
Cropland	ds		•		
water forest			0		
max_dila mean_di	tation latation	•			
max_pgv max_pgv	ја 1		•		
mean_p	JV		•	•	
() mean ep	central distance				

### Conclusions from Chile

#### 1) Static compression

(e.g., Muir-Wood and King, 1993)

2) Breaching barriers or seals (e.g., Sibson 1994; Brodsky et al. 2003)

3) Consolidation and liquefaction (e.g., Manga, 2001; Montgomery et al., 2003)

4) Shaken from the unsaturated zone (Mohr et al., 2015)

#### 5) Permeability enhancement by dynamic strain

(e.g., Rojstaczer et al., 1995; Tokunaga, 1999; Sato et al., 2000)

Supported by a variety of lab experiments and field measurements

e.g., Candela et al., *EPSL* (2014), Elkhoury et al., *JGR* (2011), Liu and Manga, *GRL* (2009) Roberts and Abdel-Fattah, *EPSL* (2009)

## Are changes only shallow? Response of springs along a fault



Response to M 5.5 Alum Rock event on October 30, 2007 King et al. (1994) documented response to 5 previous events

#### Mixture of "shallow" and "deep" water



#### Example of changes



Time since earthquake (days)

#### Features

- Flow always increases
- Peak discharge within days
- After 2 years, discharge was still elevated
- No clear temperature change
- No to very small shift towards meteoric water
- No correlation with static strains
- Clear magnitude-distance relation for response

#### Static strain?

Earthquakes followed by flow increases

Event		М	Epicentral	Volumtric	Reference
			distance	strain	
4/18/1906?	San Francisco	7.8	70 km	D	Lawson (1908)?
4/24/1984	Morgan Hill	6.2	18 km	С	King et al. (1994)
3/31/1986	Mount Lewis	5.7	15 km	С	King et al. (1994)
6/13/1988	Alum Rock	5.3	8 km	С	King et al. (1994)
4/3/1989	Alum Rock	5.0	5 km	-	King et al. (1994)
10/18/1989	Loma Prieta	7.1	40 km	D	King et al. (1994)
10/30/2007	Alum Rock	5.6	4 km	-	This study

# Springs not consistently in compressional or dilatational quadrant

Pore pressure changes  $< 10^3$  Pa

### Magnitude-distance relationship



#### Models (cartoon form)



Analytical solutions in Manga and Rowland (2009)

#### Best-fit models



Time since earthquake (days)

#### Are changes large? Widespread? Significant? Groundwater temperature in wells

Well locations with temperature data before and after the 1999 Chi-Chi earthquake



#### Taiwan



#### Pre-earthquake



#### Post-earthquake



#### Idealized problem



(b) Thermal boundary conditions





#### Permeability increases factor of 100 Returns to "normal" over a few months

Permeability is not a static quantity

<sup>1</sup>/<sub>2</sub> the transport occur over short period with enhanced permeability

### Summary

Earthquake-generated time-varying strains can change (increase) permeability for amplitudes as small at 10<sup>-6</sup>

Permeability changes in the intermediate- to farfield (a few to many fault lengths from the earthquake), and may dominate responses even in the near-field

Permeability changes can be modest, but up to 2 orders of magnitude and over regional scales

Changes persist for months to years

## Did an Earthquake Trigger the Sidoarjo ("Lusi") Mud eruption?

### What's Special About Lusi?

- Observe the birth of a large mud eruption
- Well next to main vent













#### Human and environmental effects

- ~100,000 m<sup>3</sup>/day (will last another 10-50 years?)
- >40,000 people displaced; 11 deaths
- Rice fields and shrimp ponds destroyed
- Highways, rail and power transmission disrupted
- Mud diverted to river (10 km upstream of a delta)
- Lapindo asked to pay \$280M to victims and \$140M to stop flow
- Damage estimates up to \$4.9B
- Sept 2009, House of Representatives decided event was a natural disaster

### The Yogyakarta Earthquake

- Two days before the eruption
- Mw 6.3
- 250 km away
- Earthquakes trigger other mud "volcanoes" (e.g., Pliny; Manga and Brodsky, 2006; Mellors et al., 2007; Bonini, 2009)

### Coincidence or a Trigger?

### Approach

- 1. Compare to other triggered mud volcanoes
- 2. Compare static stress changes to other stresses
- 3. Compare ground shaking to earthquakes that did not trigger an eruption
- 4. Did other earthquakes prime Lusi?
- 5. Consider the chance of fault reactivation (Mazzini et al., *EPSL* 2007)

### Response Threshold



### Conclusions

#### 1. Larger, closer earthquakes

- 2. Static stress changes too small to trigger
- 3. > 20 earthquakes caused more shaking
- 4. No unusual seismicity in months before eruption
- 5. Coulomb stresses too small and have wrong sign to reactivate fault

Nothing "special" about Yogyakarta earthquake

## Mechanism for triggering?

Use observations from Azerbaijan, Taiwan, Imperial Valley, northern Japan, Indonesia, Pakistan, Italy, Romania (Bonini et al., 2016)

•Triggered by dynamic,

not static, strains

•Repose time of ~ years

for triggering

- •More sensitive to long period waves
- •Triggering strain amplitudes too small to weaken mud or initiate undrained consolidation (our own lab experiments)
- •Bubble nucleation or growth unlikely
- •Favoured: increasing permeability and/or breaching hydraulic barrier





## A Drilling Trigger?

- Well control issues prior to the eruption
- Well was temporarily sealed
- Pressure in well, with weight of drilling mud, was enough to cause hydrofracturing or reactivate existing faults



Discussion

An alternative review of facts, coincidences and past and future studies of the Lusi eruption

Mark Tingay<sup>a</sup>, Michael Manga<sup>b, \*</sup>, Maxwell L. Rudolph<sup>c</sup>, Richard Davies<sup>d</sup>







feet metres




## Did an Earthquake Trigger the Sidoarjo ("Lusi") Mud Volcano?

N



## Did an Earthquake Trigger the Sidoarjo ("Lusi") Mud Volcano?

# Drilling trigger?

NO

## ongoing debate

Photo courtesy Arif Hidayat

## Did an Earthquake Trigger the Sidoarjo ("Lusi") Mud Volcano?

## When will it end?

NO

Photo courtesy Arif Hidayat

### **Observed Ground Deformation**



Cumulative ground deformation October 2006 - April 2011 46 L-band ALOS radar images



Shorter than our mechanical model predictions (Rudolph et al., 2011)

### Both a deep and more shallow source



Confirmed from gas geochemistry (Tingay et al. Nat Geo 2015)

#### Collaborators

Chi Wang, Joel Rowland, Emily Brodsky, Max Rudolph, Leif Karlstrom, Maria Brumm, Richard Davies, Mark Tingay, Marco Bonini, Manoo Shirzaei



Montgomery and Manga, Science (2003)

### Why?

"It's one of those curiosities of nature that has preoccupied people for years" Stuart Rojstaczer, *Nature*, 2003

- Affects water supply and water quality
  Hydrocarbon migration and recovery
  May affect underground waste repositories
- (Carrigan et al., 1991; Roeloffs, 1998)
- Affects engineering structures

New insight into

Controls on hydrologic properties Evolution of hydrologic cycle Triggered seismicity and eruptions